Lecture 6 : Analysis techniques

Decomposing the circulation

We are interested in the time- and zonal-mean circulation, the devictions from this mean, and how the two interact.

Define time mean

$$\overline{A} = \frac{1}{\varepsilon} \int_{0}^{\infty} A \, dt$$

and deviations. A' = A-Ā

Note that
$$\overline{A'} = \overline{A} - \overline{A} = \overline{A} - \overline{A} = 0$$

T is usually seasonal or longer timescale

Also define Zonal mean $\begin{bmatrix} A \end{bmatrix} = \frac{1}{2\pi} \int_{0}^{2\pi} A \, dA$ $A^{\dagger} = A - \begin{bmatrix} A \end{bmatrix}$ $\begin{bmatrix} A^{\bullet} \end{bmatrix} = 0$

Why do this?
Useful for analysing covariances
$$AB = (\overline{A} + A')(\overline{B} + B')$$
$$= \overline{A}\overline{B} + \overline{A}\overline{B}' + \overline{B}\overline{A}' + \overline{A}'\overline{B}'$$

Take the time mean:

$$\overline{AB} = (\overline{AB}' + (\overline{AB'}) * \overline{BA'} + \overline{AB'})$$

$$= \overline{AB} = \overline{AB'} = 0$$

$$\overline{AB} = \overline{AB} + \overline{A'B'}$$

$$\int_{Covariance} (1)$$

Example I

Consider the conservation equation for scalar
$$X$$

 $\frac{DX}{Dt} = 0$

Write this in flux form.

Use tangent plane equations in pressure coordinates for simplicity: $\frac{\partial X}{\partial t} + \frac{u}{2} \cdot \nabla Y = \frac{\partial X}{\partial t} + \frac{u}{\partial X} + \frac{u}{\partial Y} + \frac{u}{\partial Y} = 0 \quad (1)$

Also have mass continuity.

$$\nabla_{\rho} \cdot \underline{v} = \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial \rho} = 0 \qquad (2)$$

Add (1) \$ \$ * (2) :

$$\frac{\partial \delta}{\partial t} + \frac{\mathcal{U} \cdot \nabla_{p} \delta}{\delta t} + \delta \nabla_{p} \cdot \mathcal{U} = 0 \qquad = 0$$

$$\frac{\partial t}{\partial t} + \frac{1}{R_{e}\cos \varphi} \left[\frac{\partial (u\delta)}{\partial A} + \frac{\partial u \delta (u\delta)}{\partial \varphi} + \frac{\partial (u\delta)}{\partial \varphi} - 0 \right] = 0$$

Take zonal mean

$$\frac{\partial}{\partial t} \begin{bmatrix} x \end{bmatrix} + \frac{1}{2} \begin{bmatrix} y \end{bmatrix} = \frac{\partial}{\partial t} \begin{bmatrix} x \end{bmatrix} + \frac{\partial}{\partial t} \begin{bmatrix} y \end{bmatrix} \\ y \end{bmatrix} \\$$

The A-derivative of the divergence is 200.

$$\operatorname{Remember} \left[[\sigma X] = [\sigma][X] + [\sigma^* X^*] \right]$$

$$= 3 \qquad \frac{1}{2} \left[\frac{1}{2} + \frac{1}{2} + \frac{1}{2} \right] \left[\frac$$

Expand LHS:

$$\frac{\partial [v]}{\partial t} + \frac{[v]}{2e} \frac{\partial [v]}{\partial 4} + [v] \frac{\partial [v]}{\partial 4} = -\frac{1}{ne^{\omega s 4}} \frac{\partial}{\partial 4} \left[(v^{*} t^{*}] \cos t^{s} - \frac{\partial}{\partial t} [v^{*} t^{*}] \right]$$

$$+ \frac{[v]}{ne^{\omega s 4}} \frac{\partial [v]}{\partial 4} + \frac{[v]}{\delta p} = -\frac{1}{ne^{\omega s 4}} \frac{\partial}{\partial 4} \left[(v^{*} t^{*}] \cos t^{s} - \frac{\partial}{\partial t} [v^{*} t^{*}] \right]$$

Now, the zonal-mean continuity equation may be written, $\nabla_{r} \cdot [y] = 0$

$$=) \qquad \frac{1}{2} \quad \frac{\partial \left[u \right] \cos \phi}{\partial \phi} \quad + \quad \frac{\partial \left[u \right]}{\partial \phi} = 0$$

Hence the equation for the zonal-mean \mathcal{S} may be written: $\frac{\Im[\mathcal{S}]}{\Im\mathcal{S}} + \frac{[\mathcal{V}]}{\mathcal{N}} \frac{\Im[\mathcal{S}]}{\Im\mathcal{S}} + \frac{(\omega)\Im[\mathcal{S}]}{\Im\mathcal{S}} = -\frac{1}{\mathcal{S}} \frac{\Im[\mathcal{V}\mathcal{S}]}{\Im\mathcal{S}} - \frac{\Im}{\Im\mathcal{S}} \left[\omega^*\mathcal{S}\right]$ $\frac{\Im[\mathcal{S}]}{\mathcal{S}} + \frac{[\mathcal{U}]}{\mathcal{S}} \frac{\Im[\mathcal{S}]}{\Im\mathcal{S}} = -\nabla \cdot \left[\widetilde{\mathcal{U}}^*\mathcal{S}^*\right]$

Equation for the zonal mach:

$$\begin{aligned}
& \begin{array}{l} \sum_{x \in V} + \mu \cdot \nabla_{x} x = 0 \\
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& \end{array} \end{aligned}$$

to

Example	\mathbb{I}			
	time	g (hglng)	J (m5")	p (rgm ⁻³)
	١	0.01	3	ι
	2	0.002	-1	١
	3	0.014	6	١
		l j		

1. Calculate meridional flux of humidily

$$F(ux = P_0 \overline{Uq})$$
 {units: ms^{-1} . $kg/kg \cdot kg m^{-3}$
= $kgm^{-2}s^{-1}$

(units in p-coords are more confusing)

$$= \left[0.01(3) + 0.002(-1) + 0.04(6) \right] \frac{1}{3}$$

= 0.037 kgm²s¹

Flux by mean circulation

$$F_{mEAN} = P_0 \overline{U} \overline{q} = P_0 \left(\frac{3+6-1}{3} \right) \left(\frac{0.01+0.002+0.014}{3} \right)$$

 $= 0.023 \text{ kgm}^2 \overline{s}^{-1}$

Flux by eddies

$$F_{EODY} = P_0 \overline{u'q'} = P_0 \overline{uq} - P_0 \overline{uq} = 0.014 \text{ kgm^2s'}$$

Example III

$$[F] = P_{0}[Uq] \qquad U = AV \cos A$$

$$g = q_{0} + Dq \cos A$$

$$V_{1}(A)$$



Note that $\overline{v} = 0$, $\overline{2} = 0$

$$\begin{bmatrix} F \end{bmatrix} = p \cdot \Delta q \Delta u \end{bmatrix}^{2\pi} \cos(A) \cos(A) dA = p \cdot \Delta q \Delta u \int^{2\pi} \cos^2(A) dA > 0$$

$$\frac{1}{2\pi r} \int_{0}^{2\pi r} \cos^{2}(A) dA = \frac{1}{2\pi} \int_{0}^{2\pi r} \frac{1 + \cos(2A)}{2} dA = \frac{1}{2}$$
$$= \sum [F] = \frac{1}{2} \cos(2A)$$

Now suppose q = sin A



$$[F] = \int_{0} \underbrace{\operatorname{Av}}_{2} \int_{0}^{2\pi} \cos(A) \sin(A) dA$$

$$Let \quad u = \sin(A)$$

$$[F] = \frac{1}{2\pi} \int_{0}^{2\pi} u \, du = \frac{1}{2\pi t} \left[\frac{\sin^{2} A}{2} \right]_{0}^{2\pi} = 0$$

We know that

$$\overline{AB} = \overline{AB} + \overline{AB}'$$

Now, take the zonal mean

$$\left[\overline{AB}\right] = \left[\overline{AB}\right] + \left[\overline{AB'}\right]$$

Decompose the first term into zonal mean and eddy: $(\overline{AB}) = [\overline{A}][\overline{B}] + [\overline{A}^*\overline{B}^*] + [\overline{A}'B']$ $\overline{AB} = [\overline{A}][\overline{B}] + [\overline{A}^*\overline{B}^*] + [\overline{A}'B']$ $\overline{AB} = [\overline{A}][\overline{B}] + [\overline{A}^*\overline{B}^*] + [\overline{A}'B']$

Call also go further:

$$\left[\overline{AB}\right] = \left[\overline{A}\right]\left[\overline{B}\right] + \left(\overline{A}^{*}\overline{B}^{*}\right) + \left[A^{'}\right]\left[\overline{B}^{'}\right] + \left[\overline{A}^{*}\overline{B}^{*}\right]$$

But usually do not separate transient asymmetric from transient symmetric.

Example
Take
$$A = u$$
, $B = u$
 $\left[\overline{u^2}\right] = \left[\overline{u}\right]^2 + \left[\overline{u^{*2}}\right] + \left[\overline{u^{*2}}\right]$

Add to [J7]:

$$(u\bar{t}v^2) = (\bar{u})^2 t (\bar{v})^2 + (\bar{u}^{*2}) + (\bar{v}^{*2}) + (\bar{v}^{*2}) + (\bar{v}^{*2})$$

Inear hintic $KE of stationary $KE dt$
transient edlies
State estimation techniques
How do we calculate (\bar{u}^2) or $(\bar{u}'v')$ for
the atmosphere?
Observations are J cathered in space and time
Need a method to construct gradded estimate
Method 1: Peixoto & Cort - objective analysis
- Used to produce figures in Peixoto & Corr (1992).$

Step 1
- Ruality control data (remove outliers)
- Calculate means + cotroinies at each statia

$$\overline{T}, \overline{T'^2}, \overline{\tau T}, \overline{\tau'T'}$$

-Use atmospheric general circulation model to produce initial guess

- Correct initial guess using data assimilation techniques

Next becture: Begin to apply these techniques to understand the atmospheres angular momentum budget.