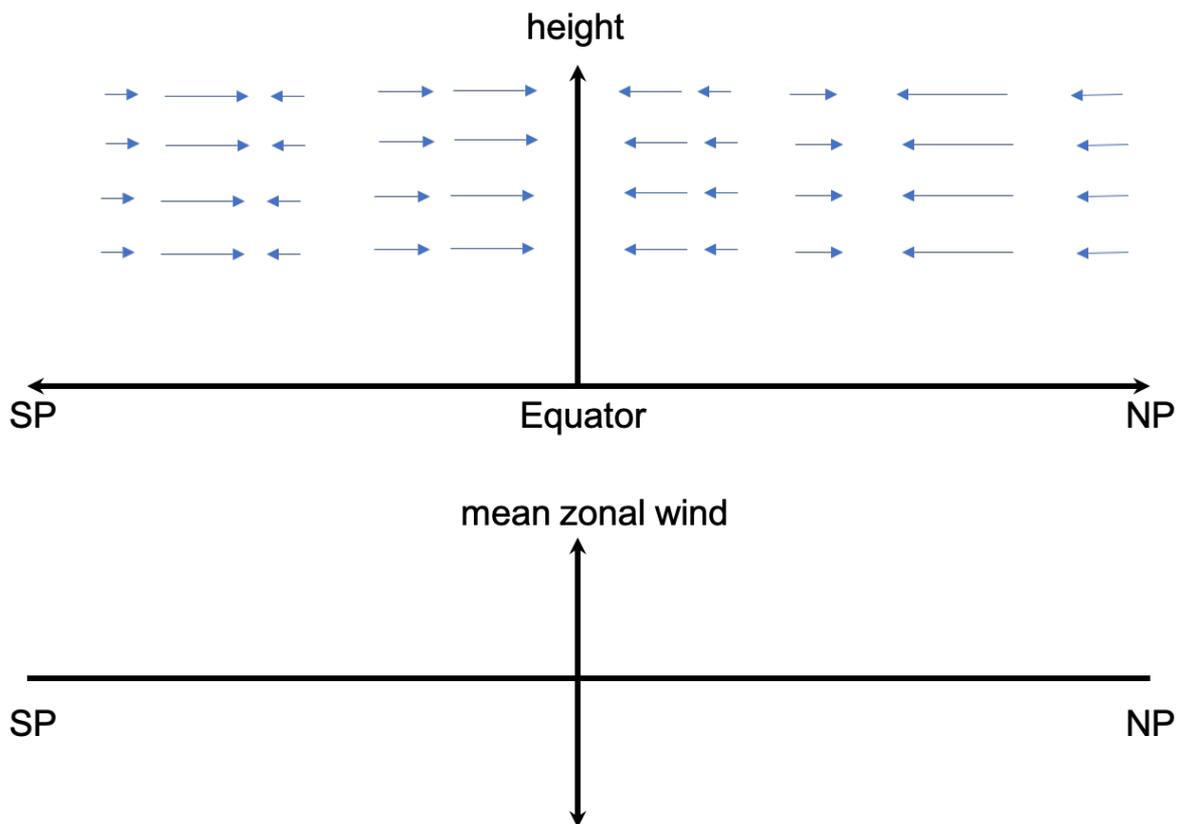


General circulation of the atmosphere Problem set 3

Problem 1.

The schematic below shows the eddy-momentum fluxes in a hypothetical atmosphere as a function of latitude (x -axis) and height (y -axis). Arrows give the value of the zonal- and time-mean eddy momentum flux at a given latitude and height up to the tropopause (arrows to the right give northward fluxes, arrows to the left give southward fluxes). Assuming quasi-geostrophic scaling is valid (except very close to the equator), sketch the corresponding overturning streamfunction (on the same axis) and the corresponding zonal-mean surface winds (on the axis below).

Arrows represent $\overline{u'v'}$, northward eddy momentum flux.



Problem 3.

Consider the distribution of zonal-mean zonal wind at 200 hPa in the figure below. Suppose a Rossby wave forms at a latitude of 60°S with a zonal phase speed of 10 m s^{-1} , and a positive meridional phase velocity. In the following, assume that the Rossby wave is barotropic, and that the relevant mean zonal background flow is the wind at 200 hPa.

- a) Assuming the zonal wavelength is “long”, so that $k \rightarrow 0$, estimate the meridional wavelength of the Rossby wave.
- b) Use linear theory (in combination with the WKB approximation) to estimate the latitude of the *critical line* at which point the wave can no longer propagate.
- c) Based on your answers to (a) and (b) above, comment on the validity of the WKB approximation for this problem.
- d) In reality, Rossby waves are not small perturbations on the zonal-mean flow. Describe what happens to a fully non-linear wave-like disturbance as it propagates meridionally from 60°S and approaches its critical line.

